

## GLOBAL MONITORING OF SPORADIC PLASMA LAYERS IN THE LOWER IONOSPHERE DURING PERIOD 2001-2008 BY USE OF CHAMP AND FORMOSAT-3 GPS OCCULTATION DATA

A. Pavelyev<sup>\*</sup>, Y. A. Liou<sup>†</sup>, J. Wickert<sup>°</sup>, N. Yen<sup>†</sup>, J. Fong<sup>†</sup> and C.C. Hsiao<sup>†</sup>

<sup>\*</sup> Kotelnikov' Institute of Radioengineering and Electronics  
Russian Academy of Sciences (IRE RAS)  
Vvedenskogo sq. 1, 141190 Fryazino, Moscow region, Russian Federation  
e-mail: pvlv@ms.ire.rssi.ru  
web page: <http://www.cplire.ru>

<sup>†</sup> S Center for Space and Remote Sensing Research, National Central University  
Chung-Li 320, Taiwan  
e-mail: yueian@csrr.ncu.edu.tw  
web page: <http://www.csrr.ncu.edu.tw>

<sup>°</sup> GeoForschungsZentrum Potsdam (GFZ-Potsdam)  
Telegrafenberg, 14473 Potsdam Germany  
E-mail: [wickert@gfz-potsdam.de](mailto:wickert@gfz-potsdam.de)  
web page <http://www.gfz-potsdam.de>

**Key words:** GPS occultation, ionosphere, plasma layers.

**Summary:** A classification of the ionospheric effect on the amplitude of radio occultation (RO) signal is introduced. Sporadic amplitude scintillation observed in RO experiments contain important information concerning the seasonal, geographical, and temporal distributions of the ionospheric disturbances and depend on solar activity. The general number of RO events with strong amplitude variations can be used as an indicator of the ionospheric activity. We found that during 2001-2008 the daily averaged  $S_4$  index measured during CHALLENGING Minisatellite Payload (CHAMP) mission depends essentially on solar activity. The maximum occurred in January 2002, minimum has been observed in summer 2008. Different temporal behavior of  $S_4$  index has been detected for polar (with latitude greater than  $60^\circ$ ) and low latitude (moderate and equatorial) regions. For polar regions  $S_4$  index is slowly decreasing with solar activity.

## 1 INTRODUCTION

The radio occultation (RO) technique is a kind of bistatic radio location when a receiver is located on board of Low Earth Orbit satellite (LEO) at an extended distance relative to transmitter of radio waves<sup>1</sup>. In distinction with the radio tomography methods (see, for example<sup>2</sup>, and references therein), RO technique may be applied practically simultaneously to investigation of both the atmosphere and ionosphere. High-stable synchronized with atomic clocks radio signals emitted at two frequencies 1575.42 MHz and 1227.6 MHz by GPS satellite navigation system are broadly used for radio occultation (RO) investigation of the Earth's plasma sheath. Plasma disturbances cause variations in the amplitude and phase of high-stable signals of navigation satellites during propagation of these signals through the Earth's magnetosphere, ionosphere, and mesosphere along the satellite-satellite paths. The amplitude scintillations are very important because in the case of layered structures they are directly connected from one side- with the eikonal acceleration and from another side- with vertical gradients of refractivity in the plasma layers<sup>3</sup>. In this contribution it is shown that (1) the  $S_4$  index of amplitude variations can be considered as an index of the ionospheric plasma influence on RO signal in the trans-ionospheric satellite-to-satellite links in a like fashion with the  $S_4$  index introduced formerly for the trans-ionospheric satellite-to-Earth links (2) the  $S_4$  index can be used in the satellite-to-satellite links as a radio-physical index of activity of plasma disturbances in the ionosphere; and (3) the relative number of GPS RO events with high values of the  $S_4$  index in the satellite-to-satellite links can be used to establish a connection between the intensity of plasma disturbances and solar activity.

## 2 VARIATIONS OF THE $S_4$ INDEX AND SOLAR ACTIVITY

The amplitude variations of RO signal can be described by the  $S_4$  scintillation index

$$s_4 = \frac{[\langle I - \langle I \rangle \rangle^2]^{1/2}}{\langle I \rangle} \quad (1)$$

where  $\langle \rangle$  is the average relevant to the height of the RO ray perigee above 40 km, and  $I(t)$  is the intensity of RO signal. When the amplitude and phase variations are caused by layered ionospheric structures, the  $S_4$  scintillation index can be connected with the refractive attenuation  $X$  and the eikonal acceleration variations  $a^3$ :

$$s_4 = [\langle X - 1 \rangle^2]^{1/2} = m[\langle a \rangle^2]^{1/2} \quad (2)$$

where  $m$  is coefficient depending on the parameters of the GPS and LEO satellites taking part in the RO experiment<sup>5</sup>. The relationships (1) and (2) give new possibility to reveal the origin of the phase and amplitude scintillations in the near-Earth propagation medium and may have a general significance for the trans-ionospheric satellite-to-satellite and satellite-to-Earth links. In the following we will indicate that the  $S_4$  scintillation index obtained from the RO amplitude data is a key parameter connecting the amplitude variations of RO signal with space weather on a global scale. Influence of sporadic solar activity can be considered in the case of solar flare during October 29–31, 2003. We treat about 2000 CHAMP RO events from October 27 to November 9, 2003. Histograms of integral distribution of the  $S_4$  index are shown in Fig. 1. A ratio  $R$  of a number of cases with the magnitude  $S_4$  larger, than value plotted on the abscissa, to the total number of measurements for the corresponding day, is plotted on the ordinates in Fig. 1. Histograms in the right panel correspond to data of  $S_4$  measurements on October 29–31, as well as on October 28, November 1, and November 4 2003, performed during relatively intense and weak disturbances in RO signal, respectively. The initial histogram values for the entire period of observations are coincident and correspond to a magnitude  $S_4$  of  $\sim 2.8\%$ , which depends on the level of fluctuations of the receiver noise. Three days, October 29–31, with considerably higher  $S_4$  levels than on the remaining days are distinguished in Fig. 1.

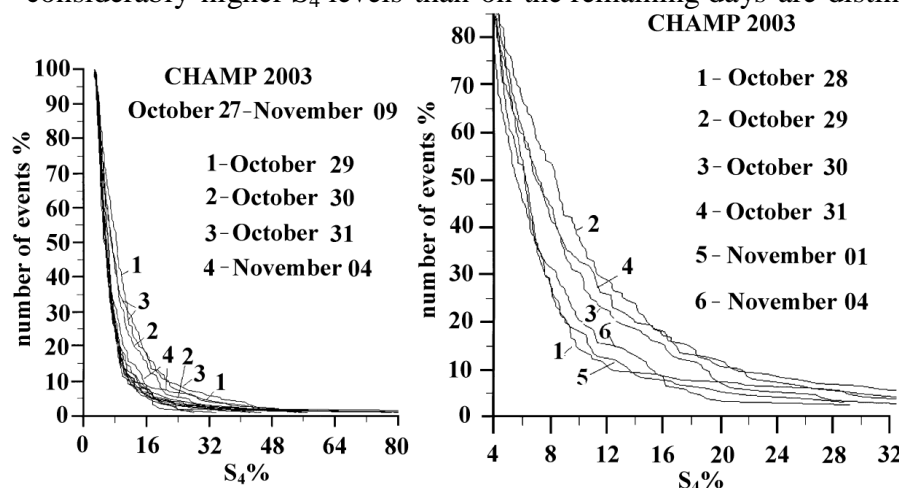


Fig. 1: Histograms of  $S_4$  index corresponding to the altitudes of CHAMP RO signal trajectory perigee higher than 40 km. Curves corresponding to the histograms for October 27 and 28 and for November 1–3 and 5–9, 2003, are almost indiscernible (left-hand panel). The right panel demonstrates the histograms in more details.

The histograms for this time interval are much higher than for the remaining days. The most difference between the histograms is observed in the 5%–19% interval of the  $S_4$  index with maximum about of 12% (Fig. 1, left and right panels). Value 12% of the  $S_4$  index is by a factor 4.3 greater than the level of the receiver's noise. Therefore, in the CHAMP RO experiments at the first

GPS frequency values of  $S_4$  index in the 5%–19% interval can be considered as the most sensitive for estimation of the ionospheric influence on RO signal. The histograms in Fig. 1 are in good agreement with variations in the hourly index of the solar wind magnetic field.

Analysis of about ~500,000 CHAMP RO session data obtained during the time period 2001-2008 allows revealing long-scale variations of the averaged index  $S_4$ . For each RO session the index  $S_4$  has been determined according to eq. (1) as an averaged value of the intensity variations when the RO ray perigee height exceed 40 km altitude. At these altitudes the ionospheric influence on the amplitude of Ro signals prevails contribution caused by the atmosphere. Therefore of about 200 values of the averaged index  $S_4$  have been obtained for different areas of the Earth per day. Then these data were averaged in the time lag 27 days and the latitude distribution of the obtained values of index  $S_4$  has been analyzed. The time dependence of  $S_{4a}$  index measured from RO CHAMP experiments during 2001-2008 is given in Fig. 2. Curve 1 corresponds to the averaged on a global scale value of the  $S_4$  index. Curve 2 indicates the averaged in the equatorial and moderate latitude areas (with latitude below  $60^\circ$ ) value of the  $S_4$  index. Curve 3 demonstrates the averaged in the polar latitudes (with latitude greater than  $60^\circ$ ) values of the  $S_4$  index. It follows from analysis of Fig. 2, that during 2001-2008 the daily globally averaged RO  $S_4$  index depends essentially on solar activity. The maximum occurred in January 2002, minimum has been observed in summer 2008. Different temporal behavior of  $S_4$  index has been detected for polar (with latitude greater than  $60^\circ$ ) and low latitude (moderate and equatorial) regions. For polar regions  $S_4$  index is slowly decreasing with solar activity. In the low latitude areas  $S_4$  index is sharply oscillating, depending on the solar ultraviolet emission variations. The different geographical behavior of  $S_4$  index indicates different origin of ionospheric plasma disturbances in polar and low latitude areas. Origin of the plasma disturbances in the polar areas may be connected with influence of solar wind, the ultraviolet emission of the Sun may be the main cause of the ionospheric irregularities in the low latitude zone. In the low latitude areas  $S_4$  index is sharply oscillating, depending on the solar ultraviolet emission variations. The different geographical behavior of  $S_{4a}$  index indicates different origin of ionospheric plasma disturbances in polar and low latitude areas. Origin of the plasma disturbances in the polar areas may be connected with influence of solar wind, the ultraviolet emission of the Sun may be the main cause of the ionospheric irregularities in the low latitude zone. Analysis of Fig. 2 reveals global oscillations with the periods of 5-7 months. Analysis of these oscillations may provide additional connection with solar activity. Therefore, the  $S_4$  index of RO signal is important radio physical indicator of solar activity.

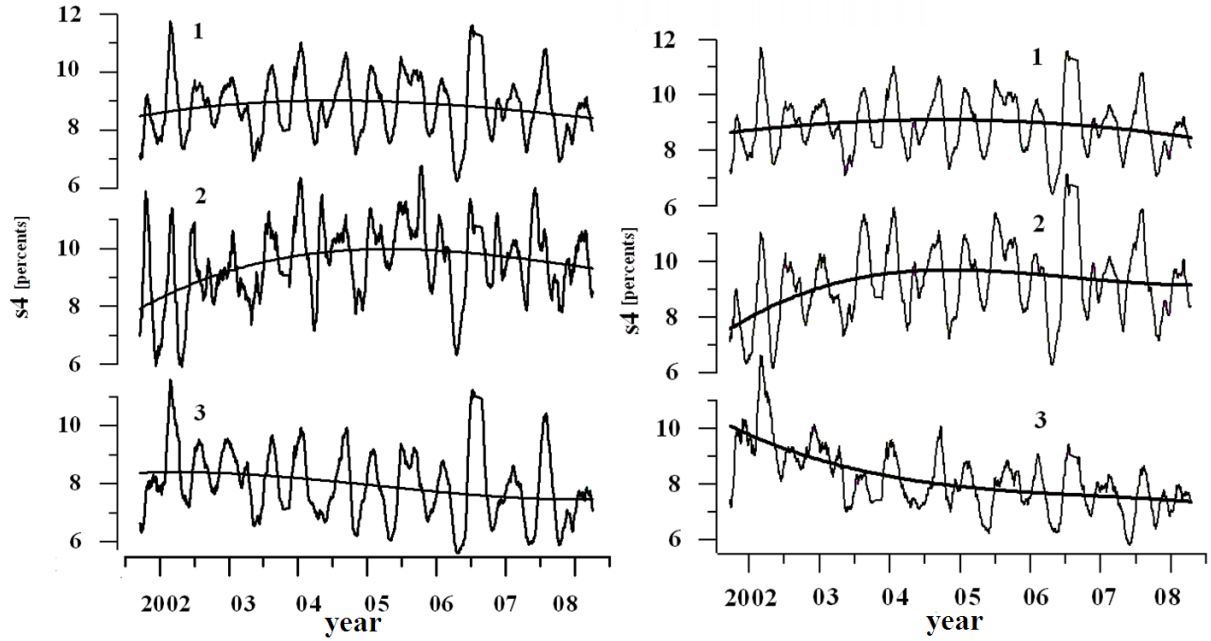


Fig. 2. Annual and latitude distribution of  $S_4$  index measured from RO CHAMP experiments. Curves 1 described the annual distribution of  $S_4$  index with a global coverage. Curves 2 correspond to the areas with latitudes below  $30^\circ$  (left) and  $55^\circ$  (right). Curves 3 correspond to the latitudes greater than  $30^\circ$  (left) and  $55^\circ$  (right). Smooth curves were obtained by a least squares method.

## REFERENCES

- [1] Fjeldbo, G., and V.R. Eshleman, The bistatic radar-occultation method for the study of planetary atmospheres, *J. Geophys. Res.*, 70(13), 3217 (1965).
- [2] Kunitsyn V.E., Tereschenko E.D. Ionospheric tomography. Springer Verlag (2003).
- [3] Pavelyev A.G., Y.A. Liou, J.Wickert, A.L. Gavrik, and C.C. Lee, Eikonal acceleration technique for studying of the Earth and planetary atmospheres by radio occultation method *Geophys. Res. Lett.* 36(21), L21807, 1-5 (2009).
- [4] Yeh, K.C., and C.H. Liu, Radio wave scintillations in the ionosphere, *Proc. IEEE*, 70(4), 324 (1982).